

23/01/12

## Fluid Dynamics ①

What is a fluid? Fluids flow ↪

Examples: Water, air, syrup, oil } Newtonian Fluids (simple)

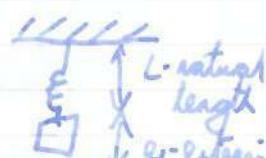
Paint, toothpaste, ketchup, soup, shampoo } Non-newtonian liquids  
Sand, salt } Granular fluids  
Foam

} Complex Fluids

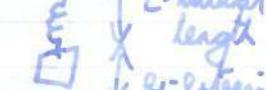
### Newtonian Fluids

These have a linear relationship between stress and rate of strain.

Stress - force per unit area (e.g. pressure)



Strain - extension per unit length (e.g. in elasticity)



Rate of Strain - rate of extension per unit length

$$\text{strain} = \frac{\delta L}{L}$$

↳ (is a gradient of velocity)

In part II, stress and strain rate are considered as tensor quantities, but here, we will consider situations that can be described not using scalar and vector fields.

We shall discuss viscous fluids. However, we shall often make an inviscid approximation.

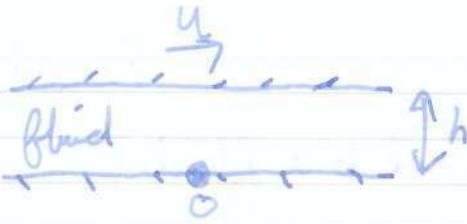
Pressure is an example of a normal stress. The pressure force per unit area on a surface with normal  $\hat{n}$ , pointing into the fluid is

$$T_p = -p \hat{n} \quad \text{surface} \quad \begin{array}{c} \hat{n} \\ \text{fluid} \\ \downarrow \\ \text{with pressure } p \end{array}$$

Gradients in pressure provide a net force.

$$\begin{array}{ccc} \text{High pressure} & \xleftarrow{\frac{\partial p}{\partial x}} & \text{Low pressure} \\ \text{body force} & \xrightarrow{-\nabla p} & \text{(force per unit volume)} \end{array}$$

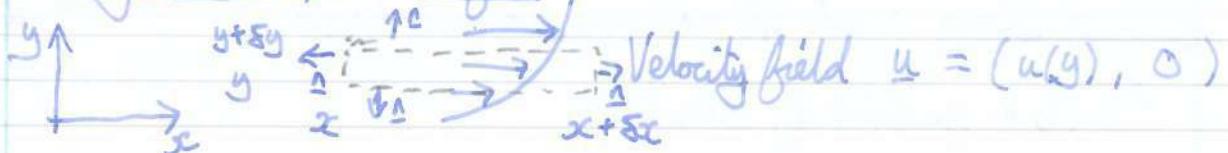
## Tangential (Shear) Stress



We find experimentally that the force per unit area required to slide plates at relative speed  $U$  is  $\tau_s \propto \frac{U}{h}$ . We write  $\tau_s = \mu \frac{U}{h}$  to define  $\mu$ , the dynamic viscosity of the fluid.

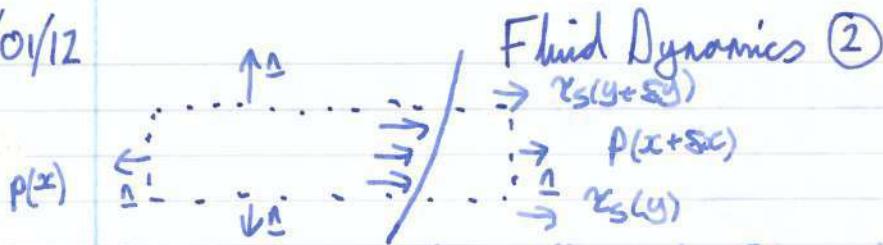
Such observations lead to the result that the tangential shear stress exerted by a fluid on a bounding surface with normal  $\mathbf{n}$  pointing into the fluid is  $\tau_s = \mu \frac{\partial u_t}{\partial n}$  where  $u_t$  is the tangential component of the fluid velocity and  $\tau_s$  is in the direction of  $u_t$ .

Steady, parallel, viscous flow,



We will start thinking about forces in the  $x$  direction, considering  $p(x)$  and  $p(x+\delta x)$ . We will also have a shear stress,  $\tau_s(y)$  and  $\tau_s(y+\delta y)$ .

29/01/12



## Fluid Dynamics ②

$\rightarrow \tau_s(y + \Delta y)$

$\rightarrow p(x + \Delta x)$

$\rightarrow \tau_s(y)$

Consider forces acting in the  $x$ -direction on the dashed slab exerted by the surrounding fluid. The slab is not accelerating so forces must balance.

$p(x) \Delta y - p(x + \Delta x) \Delta y$

$+ \tau_s(y) \Delta x + \tau_s(y + \Delta y) \Delta x = 0$

But  $\tau_s = \mu \frac{\partial u}{\partial x}$ ,  $\tau_s(y) = -\mu \frac{\partial u}{\partial y}$ ,  $\tau_s(y + \Delta y) = \mu \frac{\partial u}{\partial y}$

$- \frac{p(x + \Delta x) - p(x)}{\Delta x} + \mu \frac{u_y(y + \Delta y) - u_y(y)}{\Delta y} = 0$

$\Rightarrow - \frac{\partial p}{\partial x} + \mu \frac{\partial^2 u}{\partial y^2} = 0$

Repeat derivation in the  $y$  direction:  $- \frac{\partial p}{\partial y} = 0$

Example Sheet 1: For unsteady parallel viscous flow,

$u = (u(y, t), 0)$  with a body force (force per unit volume)

$E = (f_x, f_y)$  acting on the fluid then

$\rho \frac{\partial u}{\partial t} = - \frac{\partial p}{\partial x} + \mu \frac{\partial^2 u}{\partial y^2} + f_x \quad (\text{with } \rho \text{ the density})$

and  $O = - \frac{\partial p}{\partial y} + f_y$

Note that in a gravitational field,  $E = \rho g$

Boundary Conditions (for viscous fluids)

It has been verified experimentally (e.g. for water down to 2 molecular diameters  $\approx 6\text{\AA}$ ) that Newtonian fluids satisfy a no-slip condition, that the tangential velocity of the fluid is equal to the tangential velocity of its boundary. For a stationary, rigid boundary  $u_b = 0$ .

Stress condition: Sometimes the tangential stress  $\tau_s$  is prescribed at the boundary rather than the velocity.

If a stress  $\tau$  is applied at a boundary to the fluid then

$$-\mu \frac{\partial u}{\partial n} = \tau$$

Couette Flow - driven by boundary stresses or motions.

$$\text{Fluid } \rightarrow u(y)$$

$$\text{Stationary } \quad y=0$$

Steady, and no imposed pressure gradient, so  $\frac{\partial^2 u}{\partial y^2} = 0$ ,  $(0 < y < h)$  ①

$$\textcircled{2} \quad u = 0, (y=0), u = U, (y=h) \quad \textcircled{3} \quad \leftarrow \text{No slip}$$

$$\Rightarrow \frac{\partial u}{\partial y} = A, \quad u = Ay + B \quad \textcircled{3} \div \textcircled{B} = 0, \quad \textcircled{2} = U = Ah,$$

$$\Rightarrow u = U \frac{y}{h} \quad \text{linear velocity profile}$$

Poiseuille Flow - driven by pressure gradients

$$\rho_1 > \rho_0 \quad \rightarrow u(y) \quad \rho_0 \quad \nu_g \quad \mu \frac{\partial^2 u}{\partial y^2} = \frac{\partial p}{\partial x} \quad \textcircled{1}$$

$$0 = \frac{\partial p}{\partial y} - \rho g \quad \textcircled{2}$$

$$\text{No slip: } u=0 \quad (0 \leq y \leq h)$$

$$\textcircled{2} \Rightarrow p = \rho g y + f(x)$$

$$\textcircled{1} \Rightarrow \mu \frac{\partial^2 u}{\partial y^2} = f'(x) = -G \quad \text{constant}$$

$$\Rightarrow u = \frac{G}{2\mu} y (h-y)$$

Derived properties of a flow

1. Volume flux  $q$ : volume of fluid traversing a cross-section per unit time. For parallel flow  $q = \int_{y_1}^{y_2} u(y) dy$  per unit transverse distance

$$\text{Couette: } q = \frac{Uh}{2}, \quad \text{Poiseuille: } q = \frac{4U}{3}\nu_g$$

2. Vorticity  $\omega$ , the curl of the velocity field.  $\omega = \nabla \times \mathbf{u}$

$$\text{Couette flow: } \omega = (0, 0, -\frac{A}{h})$$

$\omega$  uniform

$\rightarrow$  rot

$$\text{Poiseuille flow: } \omega = (0, \pm \frac{2G}{3\nu_g}, 0)$$

$\omega$  parabolic

25/01/12

## Fluid Dynamics (2)

3. Surface Stress

$\tau_s$  - tangential force per unit area exerted by fluid

$$\tau_s = \mu \frac{\partial u}{\partial n}, \Delta \text{ points into the fluid}$$

Couette flow  $\tau_s = \begin{cases} \mu \frac{u}{h} & \text{for } y=0 \\ -\mu \frac{u}{h} & y=h \end{cases}$

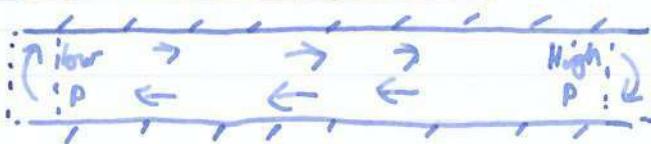
Poiseuille flow  $\tau_s = \begin{cases} \frac{Gh}{2} & \text{as } y=0 \\ 0 & y=h \end{cases}$

note that the Poiseuille flow has  $\tau_s$  independent of  $\mu$ .



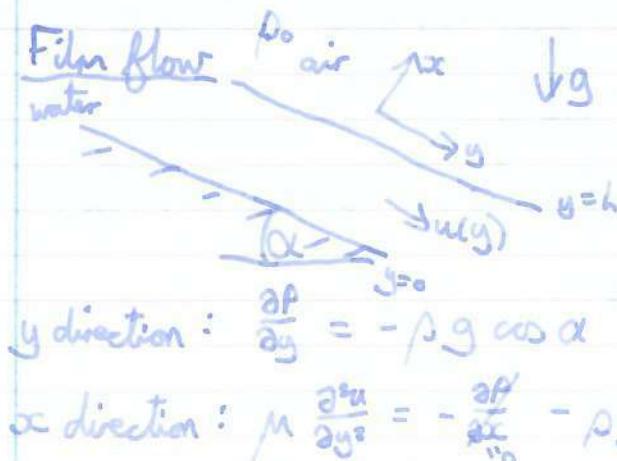
30/01/12

## Fluid Dynamics ③

Couette - Poiseuille Mixture

Pressure gradients may develop to enforce incompressibility.

No atmospheric pressure



Assume that the air exerts no tangential stress on water ( $\mu_{air} \ll \mu_{water}$ )

$$y\text{ direction: } \frac{\partial p}{\partial y} = -\rho g \cos \alpha \Rightarrow p = -\rho g \cos \alpha (y-h) + p_{atm}, \frac{\partial p}{\partial x} = 0$$

$$x\text{ direction: } \mu \frac{\partial^2 u}{\partial y^2} = -\frac{\partial p}{\partial x} - \rho g \sin \alpha$$

$$y=0, p=p_{atm}, y=h, p(x)=p_0$$

Boundary Conditions: No slip  $\Rightarrow u(x, 0) = 0$

No tangential stress  $\Rightarrow \mu \frac{\partial u}{\partial y}(0, h) = 0$

$$u(x, y) = \frac{\rho g}{2\mu} \sin \alpha y (2h-y) = \frac{g}{2\nu} \sin \alpha y (2h-y)$$

where  $\nu = \frac{\mu}{\rho}$  is the kinematic viscosity of the water.

Unsteady Parallel Viscous Flow

Consider a semi-infinite domain  $y > 0$  initially at rest. At time  $t = 0$

the boundary  $y = 0$  is set into motion with speed  $U$ .

Fluid  $\sqrt{\nu t}$

$$\rightarrow u_s(u(y, t), 0)$$

$$y\text{ direction: } \frac{\partial p}{\partial y} = \rho g, p = \rho g y + f(x)$$

Assume there is no imposed pressure gradient

$$x\text{ direction: } \rho \frac{\partial u}{\partial x} = \mu \frac{\partial^2 u}{\partial y^2} - f'(x)$$

$$u(x, t) = U \text{ for } t > 0, \quad u > 0 \text{ as } y \rightarrow \infty, \quad u(0, 0) = 0$$

$\frac{\partial u}{\partial t} = \nu \frac{\partial^2 u}{\partial y^2}$ . The velocity satisfies the diffusion equation and the kinematic viscosity  $\nu = \frac{\mu}{\rho}$  can be thought of as a diffusivity for momentum (or vorticity, see later).

The diffusion equation can be solved by Fourier transform in time (Q4), Fourier series in  $y$  (Q5, separation of variables), Laplace transforms etc.

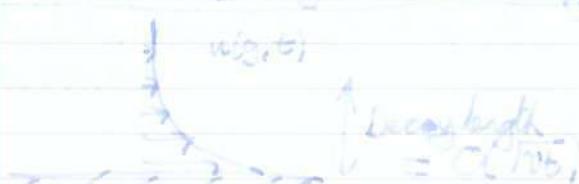
Similarity solution - see I&E Differential Equations

In the infinite domain, the diffusion equation has a similarity solution.

$$u(x, t) = U(F(\eta)), \quad \eta = \frac{x}{\sqrt{kt}}$$

$$\Rightarrow -\frac{1}{t} \eta F' = F'', \quad F = \text{erfc}(\frac{\eta}{\sqrt{2}}), \quad u = k \text{erfc}(\frac{x}{\sqrt{2kt}})$$

Kinematic vs Dynamic Viscosity



	$\mu$ (Pa s)	$\rho$ (kg/m <sup>3</sup> )	$\nu$ (m <sup>2</sup> /s)	( $-t_p$ , $20^\circ$ )
Water	$10^{-3}$	$10^3$	$10^{-6}$	$\frac{1}{20}$
Air	$2 \times 10^{-5}$	1	$2 \times 10^{-5}$	

i)  $\nu_{\text{air}} \approx 20 \nu_{\text{water}}$  so the motion is induced further into the air.

ii) Tangential Stress on  $y=0$  is  $\tau_y = \mu \frac{\partial u}{\partial y}|_{y=0} = \mu \frac{4}{2\pi} \left(\frac{2}{\pi}\right) e^{-y^2}|_{y=0}$

$$\tau_y = \mu \frac{16}{\pi} \approx 0.17$$

$$(T_0)_{\text{water}} \approx 1 \text{ (SI units)}, \quad (T_0)_{\text{air}} \approx 2.4 \times 10^{-5}$$

So water exert a much greater shear stress on the boundary than air.   
 boundary motion.   
 viscous dissipation   
 means different

Dimensional Analysis

Governing equation  $\frac{\partial u}{\partial y^2} = \frac{\partial u}{\partial t}^n$ .  $n$  - is a fracture magnitude scale of  $u$ . "T" for  $t$ , "S" - intrinsic length scale.

Note that the current problem has no extension length scales.

There is a balance between inertia and viscous dissipation when





01/02/12

## Fluid Dynamics ④

### Characterisation / Visualisation of Flow



### Lagrangian Picture

Mark (dye) a fluid particle and follow its trajectory. Trajectories can cross and it is difficult to formulate differential equations. Properties (e.g. density) would be written as  $\rho(t; \underline{x}_0)$  for a particle at time  $t$  that was released from  $\underline{x}_0$  at time  $t_0 = 0$  say.

### Eulerian Picture

Sit still and watch the world go by. Write all dependent variables (e.g.  $\underline{u}$ ) as functions of fixed location  $\underline{x}$  at time  $t$ .

### Material time derivative (Eulerian Picture)

Consider a time dependent field  $f(\underline{x}, t)$

Along a path  $\underline{x} = \underline{x}(t)$

$$\frac{df}{dt}(\underline{x}(t), t) = \frac{\partial f}{\partial x} \frac{dx}{dt} + \frac{\partial f}{\partial y} \frac{dy}{dt} + \frac{\partial f}{\partial z} \frac{dz}{dt} + \frac{\partial f}{\partial t} = \dot{\underline{x}} \cdot \nabla f + \frac{\partial f}{\partial t}$$

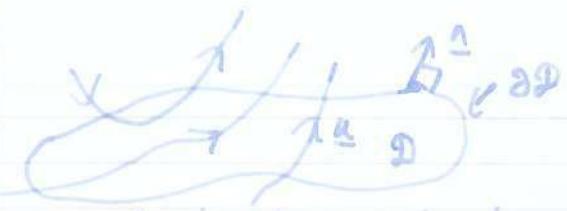


If  $\underline{x}(t)$  is the Lagrangian path followed by a fluid particle then  $\dot{\underline{x}}(t) = \underline{u}$ , the local fluid velocity. We write  $\frac{df}{dt} = \frac{Df}{Dt}$  for the time derivative along a path (trajectory) of a fluid particle.

$$\text{Material time derivative : } \frac{Df}{Dt} = \frac{\partial f}{\partial t} + \underline{u} \cdot \nabla f \leftarrow \begin{array}{l} \text{Advection time} \\ \text{derivative} \\ \text{Lagrangian time derivative} \quad \text{Eulerian time derivative} \end{array}$$

### Conservation of Mass

Consider a fixed region of space  $\mathcal{D}$  with boundary  $\partial\mathcal{D}$  and outward normal  $\mathbf{n}$ .



Mass is not created or destroyed so the mass inside  $\mathcal{D}$  can only change by a net flow across  $\partial\mathcal{D}$

$$\frac{d}{dt} \int_{\mathcal{D}} \rho dV = - \int_{\partial\mathcal{D}} \rho \mathbf{u} \cdot \mathbf{n} dS \Rightarrow \int_{\partial\mathcal{D}} \frac{\partial \rho}{\partial t} dV = - \int_{\mathcal{D}} \nabla \cdot (\rho \mathbf{u}) dV$$

using the divergence theorem.  $\Rightarrow \int_{\mathcal{D}} \frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) dV = 0$

This is true for arbitrary  $\mathcal{D}$  so  $\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$  pointwise.

This has the general form of a conservation law:

rate of change of stuff + divergence of the stuff flux = 0

Mass flux =  $\rho \mathbf{u}$

Product rule:  $\frac{\partial \rho}{\partial t} + \mathbf{u} \cdot \nabla \rho + \rho \nabla \cdot \mathbf{u} = 0$

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{u} = 0$$

If the fluid is incompressible then the density of a fluid particle cannot change  
i.e.  $\frac{D\rho}{Dt} = 0 \Rightarrow \nabla \cdot \mathbf{u} = 0$  continuity equation.

Note that in parallel flow,  $\mathbf{u} = (u, 0, 0)$   $\Rightarrow \nabla \cdot \mathbf{u} = \frac{\partial u}{\partial x} = 0$

N.B. an unconfined fluid can be treated as incompressible provided  $|u| \ll c$ , speed of sound.  $c_{air} \approx 340 \text{ ms}^{-1}$ , water  $\approx 1500 \text{ ms}^{-1}$

### Kinematic Boundary Condition

Consider a material boundary moving with velocity  $\mathbf{U}$ . 

In a local frame frame of reference moving with velocity  $\mathbf{U}$ , the fluid velocity has relative value  $\mathbf{u}' = \mathbf{u} - \mathbf{U}$  and the boundary is stationary.

  $u' = 0$  Fluid cannot cross the boundary so  $\mathbf{n} \cdot \mathbf{u}' = 0$

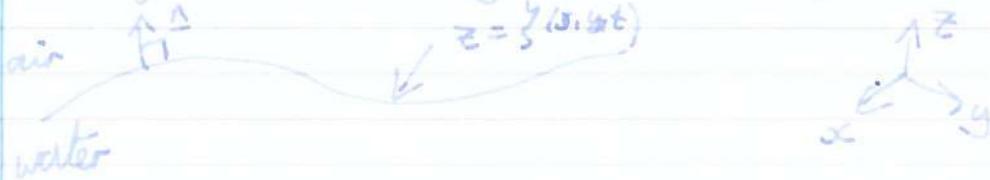
$$\Rightarrow \mathbf{n} \cdot \mathbf{u} = \mathbf{n} \cdot \mathbf{U} \text{ at boundary.}$$

01/02/12

## Fluid Dynamics ④

<sup>stationary</sup>i) At a rigid boundary,  $u = 0 \Rightarrow \nabla \cdot u = 0$ 

ii) At a free material boundary, such as the surface of a water wave,



{

Think of the surface as a contour:

$$F(x, y, z, t) = z - S(x, y, t) \quad (= 0, \text{a } 0 \text{ contour})$$

The normal to the surface is the normal to the contour of  $F$ :

$$\mathbf{n} = \nabla F = \left( -\frac{\partial S}{\partial x}, -\frac{\partial S}{\partial y}, 1 \right)$$

$$\mathbf{S} = (0, 0, \frac{\partial S}{\partial z}), \quad \mathbf{s} = (u, v, w)$$

$$\Rightarrow -u \frac{\partial S}{\partial x} = v \frac{\partial S}{\partial y} + w = \frac{\partial S}{\partial z}$$

$$\Rightarrow w = \frac{\partial S}{\partial z} + u \frac{\partial S}{\partial x} + v \frac{\partial S}{\partial y}$$

$$w = \frac{\partial S}{\partial z}$$



06/02/12

## Fluid Dynamics ⑤

### Stream functions for 2-D incompressible flow

Incompressible flow  $\Rightarrow \nabla \cdot \underline{u} = 0 \Rightarrow \underline{u} = \nabla \times \underline{A}$

for some vector potential  $\underline{A}$ .

If  $\underline{u} = (u, v, 0)$  in Cartesian components, we can find  $\underline{A} = (0, 0, \Psi(x, y))$

Note :  $u = u(x, y), v = v(x, y)$

Putting this together :  $\underline{u} = \left( \frac{\partial \Psi}{\partial y}, -\frac{\partial \Psi}{\partial x}, 0 \right)$

The scalar function  $\Psi(x, y)$  is called the stream function for the flow.

### Properties of the Stream function

The contours of  $\Psi(x, y)$  have normals  $\underline{A} = \nabla \Psi = \left( \frac{\partial \Psi}{\partial x}, \frac{\partial \Psi}{\partial y}, 0 \right)$

$$\text{So } \underline{A} \cdot \underline{u} = \frac{\partial \Psi}{\partial x} \frac{\partial u}{\partial y} - \frac{\partial \Psi}{\partial y} \frac{\partial u}{\partial x} = 0$$

So  $\underline{u}$  is orthogonal to the normals of the contours of  $\Psi$

$\Rightarrow \underline{u}$  is tangent to the contours of  $\Psi$ .

(i) The contours of  $\Psi$  are streamlines



N.B. if the flow is unsteady then streamlines are not particle paths.

The way to visualise streamlines is to seed the fluid with lots of particles and take a photograph with an open shutter of short duration

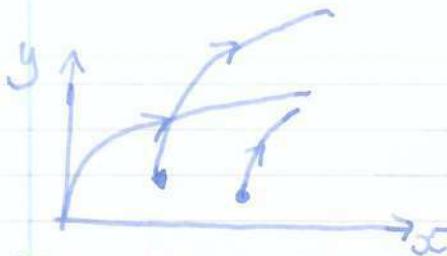
$$\text{e.g. } \underline{u} = (t, 1, 0)$$

(Lagrangian) particle paths are given

$$\begin{array}{l} \uparrow \uparrow \uparrow t=0 \\ \quad \quad \quad t=1 \end{array} \quad \text{by } \dot{x} = u, x(0) = x_0$$

$$\dot{x} = t, \dot{y} = 1, \quad x = \frac{1}{2}t^2 + x_0, \quad y = t + y_0$$

where  $t$  can be considered as a parameter for the path. Particle paths can be found by eliminating  $t \Rightarrow (x - x_0)^2 = \frac{1}{2}(y - y_0)^2$



- (ii) The flow is faster where streamlines are closer.
  - (iii) Volume flux (per unit length in the  $Z$  direction) crossing any curve from  $x_0$  to  $x_1$  =  $q = \int_{x_0}^{x_1} u \cdot n \, dL = \Psi(x_1) - \Psi(x_0)$
- 
- iv)  $\Psi$  is constant on a stationary, rigid boundary. We are always free to choose  $\Psi = 0$  on one of the stationary boundaries of a flow.

Plane-polar coordinates  $(r, \theta)$

For 2-D flow expressed in plane polar, embedded in cylindrical polar

$$(r, \theta, z) \text{ and write } \underline{u} = \nabla \times (0, 0, \Psi) = \frac{1}{r} \begin{vmatrix} \hat{e}_r & \hat{e}_\theta & \hat{e}_z \\ \frac{\partial}{\partial r} & \frac{\partial}{\partial \theta} & \frac{\partial}{\partial z} \\ 0 & 0 & \Psi \end{vmatrix}$$

$$\underline{u} = \left( \frac{1}{r} \frac{\partial \Psi}{\partial \theta}, -\frac{\partial \Psi}{\partial r}, 0 \right)$$

### Navier-Stokes Equation

Newton's 2nd law for a fluid particle is

$$\rho \frac{D\underline{u}}{Dt} = -\nabla p + \underbrace{\mu \nabla^2 \underline{u}}_{\text{mass acceleration}} + \underline{F}$$

mass acceleration = surface force + body forces

Note that the acceleration is a Lagrangian (material) derivative of the velocity.

Derivation of the viscous term  $\mu \nabla^2 \underline{u}$  requires consideration of the stress tensor, see part II fluids.

$$[\text{Note: } \nabla^2 \underline{u} = \nabla(\nabla \cdot \underline{u}) - \nabla \times (\nabla \times \underline{u})]$$

$$= -\nabla \times (\nabla \times \underline{u}) \text{ for incompressible flow}$$

In Cartesian we are straightforwardly  $\nabla^2 \underline{u} = (\nabla^2 u, \nabla^2 v, \nabla^2 w)$

06/02/12

## Fluid Dynamics ⑤

Exercise Show that if  $\underline{u} = (u_y, t, 0, 0)$  then the Navier-Stokes equations reduce to the earlier parallel flow equations derived.

In a gravitational field,  $\underline{F} = \rho \underline{g}$

Navier-Stokes equations:

$$\rho \left( \frac{\partial \underline{u}}{\partial t} + \underline{u} \cdot \nabla \underline{u} \right) = -\nabla p + \mu \nabla^2 \underline{u} + \underline{F}$$

The term  $\underline{u} \cdot \nabla \underline{u}$  is strongly non-linear and gives rise to all the interesting behaviour of fluids.



28/02/12

## Fluids, ⑥

### Hydrostatic Pressure

If  $\underline{u} \equiv 0$ , write  $p = p_H$

$$0 = -\nabla p_H + \rho g$$

$$\Rightarrow p_H = \rho g z + p_0 = p_0 - \rho g z$$



Pressure decreases upwards.

Archimedes Static force on a submerged body

Fluid density  $\rho$  

$$F = \int_{\partial F} -p_n \hat{n} dS = - \int_{\Omega} \nabla p dV$$
$$= - \int_{\Omega} \rho g dV = -\rho g \int_{\Omega} dV = -\rho g V = -M_F g$$

where  $M_F = \rho g$  is the mass of fluid displaced by the body.

Upthrust = weight of fluid displaced

Dynamic Pressure - causing or resulting from fluid flow.

Write  $p = p_H + p'$   $\approx$  dynamic pressure

$$\Rightarrow p \frac{\partial n}{\partial t} = -\nabla p' + \mu \nabla^2 \underline{u}$$

1. We usually drop the prime so that  $p$  is referred to as the dynamic pressure.
2. Every fluid particle is neutrally buoyant.
3. If there is a free surface (e.g. water/air) then we do need to consider gravity.
4. The dynamic pressure is often determined internally and is always sufficient to maintain the incompressibility constraint,  $\nabla \cdot \underline{u} = 0$

### Reynolds Number

Suppose the flow has a characteristic (typical) magnitude  $U$  and extrinsic length scale, externally imposed by geometry.

$$\rightarrow U \ll D$$



These two scales define a timescale  $T = \frac{L}{U}$

Suppose that pressure differences have characteristic magnitude  $P$ . What is the relative importance of the different terms in the Navier-Stokes equations?

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla P + 2\nu \nabla^2 \mathbf{u}$$

$$\frac{\mathbf{u}}{L} : \mathbf{u} \frac{\mathbf{u}}{L} : \frac{1}{\rho} \frac{P}{L} : 2\nu \frac{\mathbf{u}}{L^2}$$

$$\Rightarrow 1 : 1 : \frac{P}{\rho u^2} : \frac{2\nu}{uL} = \frac{1}{Re}$$

The Reynolds number  $Re = \frac{UL}{\nu}$  gives the relative magnitude of the inertial terms to the viscous terms.

N.B. Pressure must always scale to balance the dominant term in the equation so that we can impose  $\nabla \cdot \mathbf{u} = 0$

Small Re  $Re \ll 1$

Small lengths (e.g. cells), Slow flows (slow running tap)

Large viscosity (syrup, lava, oil)

Inertial terms are negligible  $P \sim \rho U^2 \frac{U}{L} = \mu \frac{U}{L}$

We can approximate by the Stokes Equations  $0 = -\nabla P + \mu \nabla^2 \mathbf{u}$

and  $\nabla \cdot \mathbf{u} = 0$ . Note that  $\mathbf{u} \propto \nabla P$

These will be studied in detail next year

Large Re  $Re \gg 1$

Viscous terms are negligible on extrinsic length scales.  $P \sim \rho U^2$ , scales with the momentum flux. On extrinsic length scales, we can approximate the NS

equations with the Euler equations:  $\rho \frac{\partial \mathbf{u}}{\partial t} = -\nabla P$ ,  $\nabla \cdot \mathbf{u} = 0$

Note that pressure gradients give rise to accelerations.

07/02/12

## Fluid Dynamics ⑥

Intrinsic length scales on which inertial terms balance viscous terms is

$$\delta \text{ where } \frac{u^2}{\delta} \sim \nu \frac{u}{\delta^2} \Rightarrow \delta \sim \frac{\nu}{u} \Rightarrow \frac{\delta}{L} = \frac{\nu}{uL} = \frac{1}{Re} \ll 1$$

At large Reynolds numbers, viscosity acts on small length scales  
(e.g. Rigid Boundaries)

$$\textcircled{1} \rightarrow u \cdot \nabla u = \textcircled{2} \textcircled{3} \begin{pmatrix} u_1 \\ u_2 \\ u_3 \end{pmatrix} \cdot \begin{pmatrix} \frac{\partial u_1}{\partial x_1} \\ \frac{\partial u_2}{\partial x_1} \\ \frac{\partial u_3}{\partial x_1} \\ \frac{\partial u_1}{\partial x_2} \\ \frac{\partial u_2}{\partial x_2} \\ \frac{\partial u_3}{\partial x_2} \\ \frac{\partial u_1}{\partial x_3} \\ \frac{\partial u_2}{\partial x_3} \\ \frac{\partial u_3}{\partial x_3} \end{pmatrix} = u_1 \frac{\partial u_1}{\partial x_1} + \dots + u_3 \frac{\partial u_3}{\partial x_3}$$

$$(u \cdot \nabla) u = \left( u_1 \frac{\partial}{\partial x_1} + \dots + u_3 \frac{\partial}{\partial x_3} \right) \begin{pmatrix} u_1 \\ u_2 \\ u_3 \end{pmatrix}$$



13/02/12

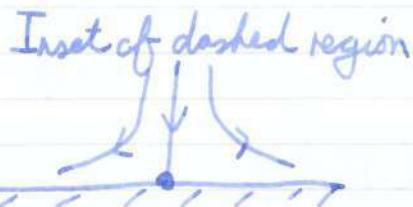
$$\rho \left( \frac{\partial \underline{u}}{\partial t} + \underline{u} \cdot \nabla \underline{u} \right) = -\nabla p + \mu \nabla^2 \underline{u}$$

Fluid Dynamics ⑦

NS

At large  $Re$ , viscous terms are unimportant on the (extrinsic) length scales of the flow.

e.g.



### A Case Study - Stagnation Point Flow

There is an exact solution of NS in which :

$$\underline{u} \sim (Ex, -Ey, 0) \text{ as } y \rightarrow \infty, \quad \underline{u} = 0 \text{ on } y = 0$$

No slip

$$\text{The solution has the form } \underline{u} = (Ex g'(\eta), -\sqrt{vE} g(\eta), 0)$$

$$\text{where } \eta = \frac{y}{\delta} = \sqrt{\frac{v}{E}} y. \quad \text{The (intrinsic) length scale } \delta = \sqrt{\frac{v}{E}}$$

### Exercises

i) Show that  $\nabla \cdot \underline{u} = 0$  and that the stream function

$$\Psi = \sqrt{vE} \xi g(\eta) \text{ with } \Psi \sim Exy \text{ as } y \rightarrow \infty$$

ii) Show that the NS equations give

$$Ex(g'^2 - gg'') = -\frac{1}{\rho} p_{x0} + E^2 x g''' \quad ①$$

$$E\sqrt{vE} gg' = -\frac{1}{\rho} p_y - E\sqrt{vE} g'' \quad ②$$

Differentiate :

$$\frac{\partial}{\partial x} ② \Rightarrow p_{x0} = 0$$

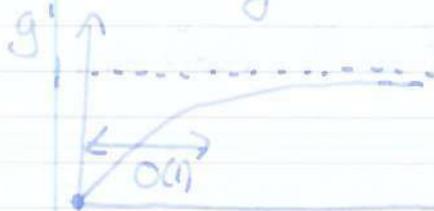
$$\frac{\partial}{\partial y} ① \Rightarrow (g'^2 - gg'')' = g^{(4)}$$

iii) Show that the boundary conditions give

$$g' \geq 1, \quad g \sim \eta \text{ as } \eta \rightarrow \infty, \quad g = g' = 0 \text{ when } \eta = 0$$

All dimensional parameters have been absorbed into the scaled variable,  $g$  and  $\eta$ .

So we only need to solve the ODE once to find



The far-field velocity  $\underline{u} = (E_x, -E_y, 0)$

is reached to an excellent approximation

$$\text{when } \theta \gg 1 \Rightarrow y \approx \delta = \sqrt{\frac{E}{\nu}}$$

### Overall Picture



Recall that  $\underline{u} = (E_x, -E_y, 0)$  in the far field.

If we're interested in the flow on scales much larger than  $\delta$  then solve the equations in  $y \geq \delta$ , ignoring the viscous terms.

Formally, we will take the limit  $\delta \rightarrow 0$  and solve the Euler equations in  $y > 0$ . BUT we can't then use the no-slip condition.

### Inviscid Approximation

If  $Re \gg 1$  then solve the Euler equations

$$\rho \frac{D\underline{u}}{Dt} = -\nabla p \quad (+\ddot{x})$$

$$\nabla \cdot \underline{u} = 0, \quad \underline{u} \cdot \underline{n} = 0 \text{ at stationary, rigid boundaries.}$$

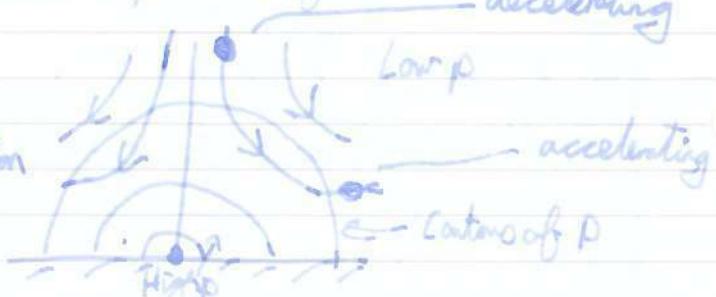
The no-slip condition cannot be imposed on solutions of the Euler equation.

Exercise Show that the flow

$\underline{u} = (E_x, -E_y, 0)$  satisfies the Euler equations in  $y > 0$  with a rigid boundary at  $y = 0$  and a pressure field  $p = p_0 - \frac{\rho E^2}{2} (x^2 + y^2)$

$$p = p_0 - \frac{\rho E^2}{2} (x^2 + y^2)$$

The pressure field acts as an internal reaction force imposing the constraint of incompressibility.



15/02/12

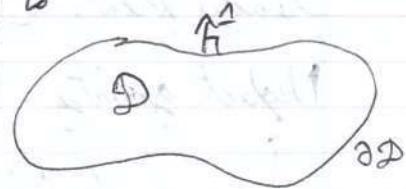
# Fluid Dynamics ②

## Momentum Equation

For incompressible, inviscid flow. The momentum of fluid inside a region  $D$  with boundary  $\partial D$  can change owing to



i) Momentum flowing across the boundary.



ii) Volume (body) forces

iii) Surface pressure forces

iv) Surface viscous forces : tangential shear stress, normal viscous stress

$$(*) \frac{d}{dt} \int_D \rho \underline{u} dV = - \int_{\partial D} \rho \underline{u} \underline{n}_{(i)} \cdot \underline{n} dS + \int_D \underline{f} dV - \int_{\partial D} P \underline{n}_{(ii)} dS$$

In components :

$$\frac{d}{dt} \int_D \rho u_i dV = - \int_{\partial D} \rho u_i u_j n_j dS + \int_D f_i dV - \int_{\partial D} P n_i dS$$

$$\Rightarrow \int_D \rho \frac{\partial u_i}{\partial t} + \rho \frac{\partial}{\partial x_j} (u_i u_j) dV = \int_D - \frac{\partial P}{\partial x_i} + f_i dV$$

This is true for arbitrary regions  $D$ .

$$\rho \frac{\partial u_i}{\partial t} + \rho \left( u_j \frac{\partial u_i}{\partial x_j} + u_i \frac{\partial u_j}{\partial x_j} \right) = - \frac{\partial P}{\partial x_i} + f_i$$

$$\rho \frac{\partial u_i}{\partial t} + \rho (\underline{u} \cdot \nabla) \underline{u} + \underbrace{(\nabla \cdot \underline{u}) \underline{u}}_{=0} = - \nabla P + \underline{f}$$

$$\Rightarrow \rho \frac{D \underline{u}}{Dt} = - \nabla P + \underline{f} \quad \text{Euler Momentum Equation}$$

$\frac{Du}{Dt}$  is the acceleration of a fluid particle,  $\frac{D \underline{u}}{Dt}$  is the local rate of change of the velocity at a point

## Conservative body forces

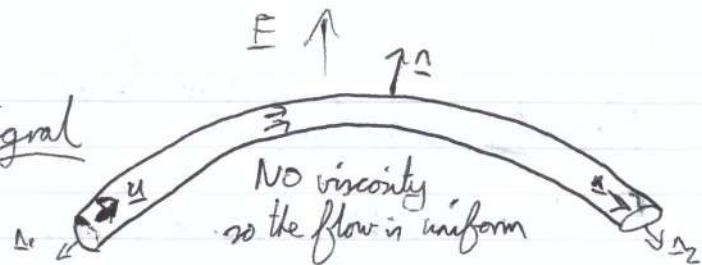
$$\underline{F} = - \nabla X \quad \text{e.g. gravity, } \underline{F} = \rho \underline{g} = \nabla (\rho g z), X = -\rho g z$$

## Momentum Integral for steady flow

$$\begin{aligned} \text{From } (*) , \quad 0 &= - \int_{\partial D} \rho \underline{u} (\underline{u} \cdot \underline{n}) dS - \int_D \nabla X dV - \int_{\partial D} P \underline{n} dS \\ \Rightarrow \int_{\partial D} \rho \underline{u} (\underline{u} \cdot \underline{n}) + P \underline{n} + X \underline{n} dS &= 0 \end{aligned}$$

## Application of Momentum Integral

Curved hose-pipe



Steady flow. Conservation of mass  $\Rightarrow |u| = \text{constant} = u$

Neglect gravity. No acceleration along the pipe.  $\Rightarrow p$  is constant

$$\int_{\text{curved surface}} \rho u (\underline{n}) + p \underline{n} dS + A_p (-u \underline{n}_1) (-u) + A_p u \underline{n}_2 u = 0$$

Cross sectional Area  $\rightarrow$

$$+ A_p \underline{n}_1 + A_p \underline{n}_2$$

Force on pipe :

$$F = \int p \underline{n} dS = -A(\rho u^2 + p)(n_1 + n_2)$$

## Bernoulli's Equation

$$\text{Vector identity: } \underline{u} \times (\nabla \times \underline{u}) = \nabla \left( \frac{1}{2} \rho |\underline{u}|^2 \right) - (\underline{u} \cdot \nabla) \underline{u}$$

$$\text{From Euler's Equation } \rho \frac{\partial \underline{u}}{\partial t} + \nabla \left( \frac{1}{2} \rho |\underline{u}|^2 \right) - \rho \underline{u} \times (\nabla \times \underline{u}) = -\nabla p - \nabla \chi$$

(\*\*\*)

Dot with  $\underline{u}$  :

$$\textcircled{1} \text{ Steady Flow } \underbrace{\underline{u} \cdot \nabla}_{\perp} \left( \frac{1}{2} \rho |\underline{u}|^2 + p + \chi \right) = 0$$

Directional derivative with respect to time along a streamline

$$\Rightarrow H = \frac{1}{2} \rho |\underline{u}|^2 + p + \chi = \text{constant along streamline}$$

Bernoulli's Theorem.

Use the divergence theorem recalling  $\nabla \cdot \underline{u} = 0$

$$\textcircled{2} \text{ Integrate } (*) \quad \frac{d}{dt} \int_D \frac{1}{2} \rho |\underline{u}|^2 dV = - \int_{\partial D} H \underline{u} \cdot \underline{n} dS$$

So the rate of change of the kinetic energy inside  $D$  is equal to the flux of  $H$  across the boundary of  $D$ . So  $H$  is the transportable energy of the fluid.

## Applications of Bernoulli's Equation

i) Pitot tube for measuring air speed.



Along stagnation point streamline

Stagnation Point

$$\frac{1}{2} \rho u^2 + p_0 = p_1 \quad \Rightarrow \quad u = \sqrt{2 \frac{p_1 - p_0}{\rho}}$$

Measuring  $p_1$  gives  $u$ . In the barometer  $\frac{p_1 - p_0}{\rho} = \rho_w g h$

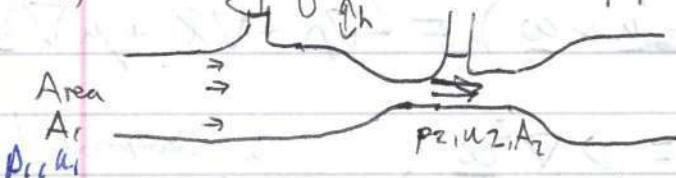
20/02/12

## Fluid Dynamics ⑨

Bernoulli's Theorem

$$H \equiv \frac{1}{2} \rho u^2 + p + \chi = \text{constant along streamlines}$$

ii) Measuring flow rate in a pipe

Conservation of mass : Volume flow rate  $q = A_1 u_1 = A_2 u_2$ 

Bernoulli :  $H_1 = H_2$

$$\frac{1}{2} \rho u_1^2 + p_1 = \frac{1}{2} \rho u_2^2 + p_2$$

$$\frac{1}{2} \rho \left( \frac{q}{A_1} \right)^2 + p_1 = \frac{1}{2} \rho \left( \frac{q}{A_2} \right)^2 + p_2$$

$$\Rightarrow \left( \frac{1}{A_2^2} - \frac{1}{A_1^2} \right) q^2 = \frac{2}{\rho} (p_1 - p_2) = \frac{2 \rho g h}{\rho} = 2 g h$$

$$\Rightarrow q = \sqrt{2 g h} \frac{A_1 A_2}{|A_1^2 - A_2^2|}$$

Linear flowsConsider the flow in the neighbourhood of a fixed point  $\underline{x}_0$ 

$$\underline{u}(\underline{x}) = \underline{u}(\underline{x}_0) + (\underline{x} - \underline{x}_0) \cdot \nabla \underline{u}(\underline{x}_0) + \dots$$

$$\approx \underline{u}_0 + \underline{\epsilon} \cdot \nabla \underline{u} \quad \text{- linear approximation, } \underline{\epsilon} = \underline{x} - \underline{x}_0$$

where  $\underline{u}(\underline{x}_0) = \underline{u}_0$  is a constant vector

$$\nabla \underline{u} = \frac{\partial \underline{u}_i}{\partial x_j} = E_{ij} + \Omega_{ij} = \underline{E} + \underline{\Omega}$$

$$E_{ij} = \frac{1}{2} \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right), \text{symmetric.} \quad \Omega_{ij} = \frac{1}{2} \left( \frac{\partial u_i}{\partial x_j} - \frac{\partial u_j}{\partial x_i} \right), \text{antisymmetric}$$

Recall: vorticity  $\underline{\omega} = \nabla \times \underline{u}$ 

$$\text{Note: } \underline{\omega} \times \underline{\epsilon} = (\nabla \times \underline{u}) \times \underline{\epsilon} = \left( \frac{\partial u_i}{\partial x_j} - \frac{\partial u_j}{\partial x_i} \right) r_i = 2 \Omega_{ij} r_i = 2 \underline{\Omega} \cdot \underline{\epsilon}$$

Therefore

$$\underline{u} \approx \underline{u}_0 + \underline{E} \cdot \underline{\epsilon} + \frac{1}{2} \underline{\omega} \times \underline{\epsilon}$$



Uniform Flow



Pure Strain



Pure Rotation

Note that the local rotation rate =  $\frac{1}{2}$  the vorticityNote also that the Strain Rate tensor  $\underline{\epsilon}$  is symmetric, and hence diagonalisable with orthogonal eigenvectors, and  $\underline{\epsilon}$  is traceless, because  $\nabla \cdot \underline{u} = 0$ 

$$E_{kk} = \frac{\partial u_k}{\partial x_k} = \nabla \cdot \underline{u} = 0$$

With respect to principle axes,  $\underline{\epsilon} = \begin{pmatrix} E_1 & & \\ & E_2 & \\ & & E_3 \end{pmatrix}$  where  $E_1 + E_2 + E_3 = \text{real}$

## Vorticity Equation

The Navier-Stokes equations for a viscous fluid gives

$$\rho \left( \frac{\partial \underline{u}}{\partial t} + \underline{u} \cdot \nabla \underline{u} \right) = -\nabla p - \nabla X + \mu \nabla^2 \underline{u}$$

$$\Rightarrow \rho \left( \frac{\partial \underline{u}}{\partial t} + \nabla \left( \frac{1}{2} (\underline{u})^2 \right) - \underline{u} \times \underline{\omega} \right) = -\nabla p - \nabla X + \mu \nabla^2 \underline{u}$$

Take the curl:

$$\frac{\partial \underline{\omega}}{\partial t} - \nabla \times (\underline{u} \times \underline{\omega}) = \nu \nabla^2 \underline{\omega} \quad \text{○ incompressibility}$$

$$\text{But } \nabla \times (\underline{u} \times \underline{\omega}) = (\nabla \cdot \underline{\omega}) \underline{u} + (\underline{\omega} \cdot \nabla) \underline{u} - (\nabla \cdot \underline{u}) \underline{\omega} - (\underline{u} \cdot \nabla) \underline{\omega}$$

$$\text{So } \frac{\partial \underline{\omega}}{\partial t} + \underline{u} \cdot \nabla \underline{\omega} = \underline{\omega} \cdot \nabla \underline{u} + \nu \nabla^2 \underline{\omega}$$

$$\frac{\partial \underline{\omega}}{\partial t} = \underline{\omega} \cdot \nabla \underline{u} + \nu \nabla^2 \underline{\omega}$$

$\frac{\partial \underline{\omega}}{\partial t}$  is the local rate of change of vorticity

$\underline{u} \cdot \nabla \underline{\omega}$  represents advection of vorticity

$\underline{\omega} \cdot \nabla \underline{u}$  represents amplification (or reduction) of vorticity by stretching (or compression)

$\nu \nabla^2 \underline{\omega}$  represents dissipation of vorticity by the action of viscosity, and also allows for the generation of vorticity by the no-slip condition at rigid boundaries.

## Vortex amplification by stretching

The vorticity equation in an inviscid fluid is

$$\underline{\omega} \cdot \frac{D \underline{\omega}}{Dt} = \underline{\omega} \cdot \nabla \underline{u} \cdot \underline{\omega}$$

$$\Rightarrow \frac{D}{Dt} \left( \frac{1}{2} \underline{\omega}^2 \right) = \underline{\omega} \cdot \left( E + \frac{\Omega}{\underline{\omega}} \right) \cdot \underline{\omega} = \underline{\omega} \cdot E \cdot \underline{\omega}$$

w.r.t principal axes of  $E$

$$\frac{D}{Dt} \left( \frac{1}{2} \underline{\omega}^2 \right) = E_1 \underline{\omega}_1^2 + E_2 \underline{\omega}_2^2 + E_3 \underline{\omega}_3^2$$

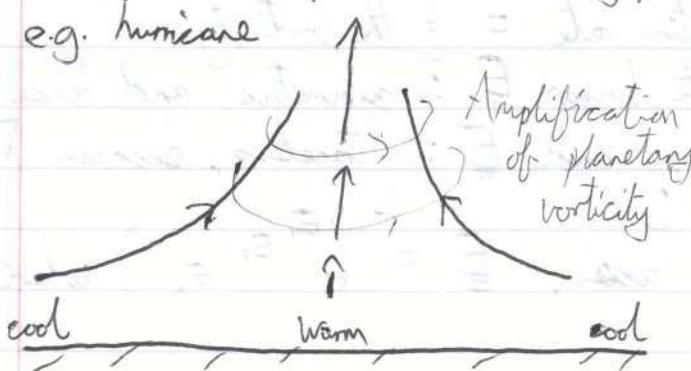
Suppose  $\underline{\omega} = (\underline{\omega}_1, 0, 0)$ .

Then  $\frac{D}{Dt} \left( \frac{1}{2} \underline{\omega}_1^2 \right) = E_1 \underline{\omega}_1^2$  and the vorticity grows or decays exponentially depending on the sign of  $E_1$ .

Note that the flow associated with  $E$  is  $\underline{u} = (E_{1x}, E_{2y}, E_{3z})$

So  $E_1 > 0$  corresponds to stretching parallel to the  $x$  axis.

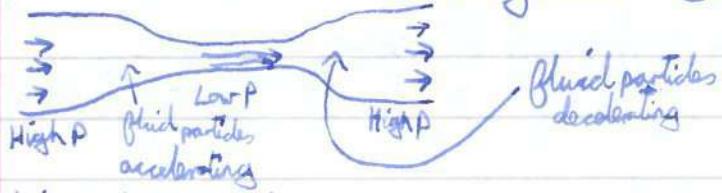
e.g. hurricane



The hurricane circulates cyclonically  
(in the direction of Earth's rotation)

22/02/12

## Fluid Dynamics ⑩



### Vorticity Equation

$$\frac{\partial \underline{\omega}}{\partial t} = \underline{\omega} \cdot \nabla \underline{u} + 2 \nabla^2 \underline{u}$$

not in this course

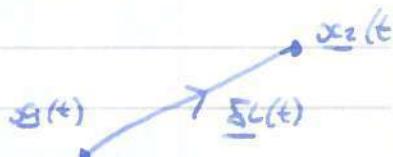
### Vortex Stretching

Considering two neighbouring (Lagrangian) fluid particles.

$$\frac{D \underline{x}_2}{D t} = \underline{u}(\underline{x}_2), \quad \frac{D \underline{x}_1}{D t} = \underline{u}(\underline{x}_1)$$

$$\Rightarrow \frac{D}{D t} \underline{\delta L} = \underline{u}(\underline{x}_2) - \underline{u}(\underline{x}_1) = \underline{u}(\underline{x}_1) + \underline{\delta L} \cdot \nabla \underline{u}(\underline{x}_1) + \dots - \underline{u}(\underline{x}_1) \\ = \underline{\delta L} \cdot \nabla \underline{u} \text{ to leading order}$$

$\underline{\delta L}$  satisfies the same equation as  $\underline{\omega}$ , so vorticity is stretched as fluid elements are stretched.



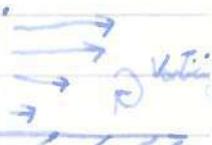
### Inviscid (homogeneous) fluid

$$\frac{\partial \underline{\omega}}{\partial t} = \underline{\omega} \cdot \nabla \underline{u} - \underline{u} \cdot \nabla \underline{\omega}$$

Vorticity can be amplified, and advected, but cannot be generated.

#### Airfoil

i) Vorticity is generated by viscous action near rigid boundaries.



ii) Vorticity can also be generated by buoyancy

✓ Inviscid rules this out



✓ Homogeneity rules this out

### Inviscid, irrotational flow

If  $\nabla \times \underline{u} = 0$  at  $t = 0$ , and the fluid is inviscid, then

$\nabla \times \underline{u} = 0$  for all time

$\Rightarrow \underline{u} = \nabla \varphi$  for some scalar "velocity potential"

Incompressibility :  $\nabla \cdot \underline{u} = 0 \Rightarrow \nabla^2 \varphi = 0$ , Laplace's Equation

### Three-dimensional flows

i) Spherically symmetric flow  $\frac{1}{r^2} \frac{\partial}{\partial r} (r^2 \frac{\partial \varphi}{\partial r}) = 0$

$$\frac{\partial \varphi}{\partial r} = \frac{A}{r^2} \Rightarrow \varphi = -\frac{A}{r} + B \quad (B=0 \text{ wlog})$$

$$\text{Velocity field } \underline{u} = \nabla \varphi = \frac{\partial \varphi}{\partial r} \underline{e}_r = \frac{A}{r^2} \underline{e}_r$$

Volume flux across the surface of the sphere,  $r=a$  is  $\frac{q}{4\pi r^2} 4\pi a^2 = q$  independent of  $A$

$$q_s = \int_S \frac{1}{r} \cdot \hat{n} dS = \int_S \frac{\hat{A}}{a} \cdot \hat{n} dS = \int_S \frac{\hat{A}}{a^2} dS = \frac{4}{a^2} 4\pi a^2 = 4\pi A$$

$$\Rightarrow q = -\frac{q_s}{4\pi r^2}$$

Represents a point source of fluid of strength (volume flux)  $q$ , at the origin.

Note:  $\nabla^2 \phi = q_s \delta(\mathbf{r})$ , so  $\phi$  is the 3D Green's function for potential flow.

### i) Axisymmetric Solutions

In spherical polar coordinates,  $\nabla^2 \phi = \frac{1}{r^2} \frac{\partial^2}{\partial r^2} (r^2 \frac{\partial \phi}{\partial r}) + \frac{1}{r^2 \sin \theta} \frac{\partial^2}{\partial \theta^2} (\sin \theta \frac{\partial \phi}{\partial \theta})$

$$\Rightarrow \phi = \sum_{n=0}^{\infty} (A_n r^n + B_n r^{-n-1}) P_n(\cos \theta) \text{ with } P_n \text{ the } n^{\text{th}} \text{ Legendre polynomial}$$

e.g. uniform flow past a sphere.

$$u_\infty \rightarrow \begin{array}{c} \circlearrowleft \\ \rightarrow \\ \rightarrow \end{array} T^\infty \rightarrow$$

In Cartesian coordinates,  $\underline{u}_\infty = (0, 0, U)$   
 $u_\infty = U \hat{z} = U r \cos \theta$

- ①  $\nabla^2 \phi = 0 \quad (r > a)$
- ②  $\phi \sim U r \cos \theta \text{ as } (r \rightarrow \infty)$
- ③  $u \cdot \hat{n} = \frac{\partial \phi}{\partial r} = 0 \text{ at } r=a \text{ (fluid cannot penetrate a rigid boundary)}$

Notes: equation is linear.

forcing  $\propto \cos \theta$  ( $= P_1(\cos \theta)$ )

$P_1(\cos \theta)$  is an eigenfunction of  $\nabla^2$  in spherical polar.

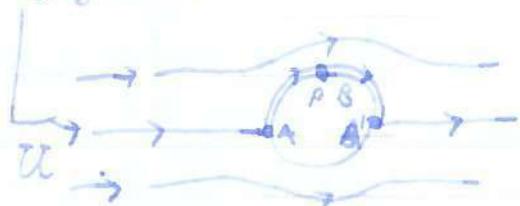
So the solution is  $\phi = f(r) P_1(\cos \theta) = (Ar + \frac{B}{r^2}) P_1(\cos \theta)$

$$\text{② } \Rightarrow A = U \quad \text{③ } \Rightarrow A - \frac{B}{a^3} = 0,$$

$$\phi = U \left( r + \frac{a^3}{2r^2} \right) \cos \theta$$

uniform flow  $\rightarrow$  dipole field

Stagnation Point Streamline



### Velocity and Pressure

$$u_r = \frac{\partial \phi}{\partial r} = U \left( 1 - \frac{a^3}{2r^2} \right) \cos \theta$$

$$u_\theta = \frac{1}{r} \frac{\partial \phi}{\partial \theta} = -U \left( 1 + \frac{a^3}{2r^2} \right) \sin \theta$$

Note that  $u_\theta \neq 0$  on  $r=a$ .

Use Bernoulli's Theorem on the stagnation point streamline

$$\frac{1}{2} \rho U^2 + p_\infty = p + \frac{1}{2} \rho u^2 \frac{9}{4} \sin^2 \theta, \quad p = p_\infty + \frac{1}{2} \rho U^2 \left( 1 - \frac{9}{4} \sin^2 \theta \right)$$

Note:  $p_A = p_{A'} = p_\infty + \frac{1}{2} \rho U^2$  (high)

$$p_B = p_\infty - \frac{3}{4} \rho U^2 \quad (\text{low})$$

Pressure is symmetric fore and aft (around the equator), so the net force on the sphere is zero! (D'Alembert's Paradox)

27/02/12

## Fluid Dynamics (II)

Solid Sphere (non-examinable)

Empirically,  $F = \rho U^2 \pi a^2 \times \frac{1}{2} C_D$ , where  $C_D$  is a measured drag coefficient.  
 $C_D = C_D(Re) \approx 0.4$  for large  $Re$

Bubble

Potential flow solution is reasonable in several circumstances.

$$\text{Exercise: K.E. of fluid} = \int_{\text{bubble}} \frac{1}{2} \rho u^2 dV = \frac{\pi}{3} a^3 \rho U^2$$

$$= \frac{1}{2} M_A U^2$$

where  $M_A$  is the "added mass" (better terminology might be "added inertia")

$$M_A = \frac{1}{2} \left( \frac{4}{3} \pi a^3 \right) \rho = \frac{1}{2} M_0 \quad \text{where } M_0 \text{ is the mass of fluid displaced.}$$

$$\text{Change in P.E.} = - M_0 g h \quad (\text{missing mass})$$

$$\text{So } \frac{1}{2} M_A U^2 - M_0 g h = E \text{ (constant)}$$

$$\Rightarrow M_A U \dot{U} - M_0 g \dot{h} = 0 \quad (\dot{h} = \dot{U})$$

$\Rightarrow \dot{U} = 2g$ , so the bubble accelerates (upwards) at  $2g$ .

Two dimensional Potential Flows

$$(i) \text{ Point Source: } \nabla^2 \phi = \frac{1}{r} \frac{\partial^2}{\partial r^2} \left( r \frac{\partial \phi}{\partial r} \right) = 0 \text{ for anti-symmetric flow}$$

or From mass conservation:

$$\frac{\partial \phi}{\partial r} = \frac{Q}{r^2} \quad \Rightarrow \frac{\partial u}{\partial r} = \frac{Q}{r^2} \quad \Rightarrow u = \frac{Q}{2\pi r} + C_1 \quad \phi = -\frac{Q}{2\pi} \ln r + C_2$$

ii) General Solution in Plane Polar coordinates

$$\phi = \frac{k}{2\pi} \ln r + r \frac{B_0}{2\pi} \theta + \sum_{n=1}^{\infty} (A_n r^n + B_n r^{-n}) \left[ \begin{array}{l} \text{for } r > 0 \\ \text{for } r < 0 \end{array} \right]$$

$$(iii) \text{ Point Vortex: } \phi = \frac{k}{2\pi} \theta \Rightarrow u = \nabla \phi = \frac{1}{r} \frac{\partial \phi}{\partial r} e_r = \frac{k}{2\pi r} e_\theta$$

Exercise: Show that  $\nabla \times u = 0$  if  $r \neq 0$

and that  $\oint_C u \cdot d\ell = \begin{cases} k & \text{if the origin is inside } C \\ 0 & \text{otherwise} \end{cases}$

The circulation  $\oint_C u \cdot d\ell = \int_S \omega \cdot n dS \quad [\omega = \nabla \times u]$  is a measure of the vorticity enclosed within  $C$ .

(iv) Uniform Flow past a circle (or cylinder)

$$\nabla^2 \phi = 0 \quad (r > a)$$

$$\phi \approx U r \cos \theta \quad (r \rightarrow \infty)$$

$$u_r = \frac{\partial \phi}{\partial r} = 0 \quad (r=a), \quad \phi = U(r + \frac{a^2}{r}) \cos \theta + \frac{k}{2\pi} \theta$$

No net source of fluid, so  $q_r = 0$  but we must allow a non-zero  $k$  to account for any vorticity in the viscous boundary layer near the surface of the circle.

$$\text{Velocity } u_r = \frac{\partial \phi}{\partial r} = U(1 - \frac{a^2}{r^2}) \cos \theta, \quad u_\theta = \frac{1}{r} \frac{\partial \phi}{\partial \theta} = -U(1 + \frac{a^2}{r^2}) \sin \theta + \frac{k}{2\pi r}$$

Streamfunction (c.f Sheet 1, Q8)

$$\psi = U r \sin \theta (1 - \frac{a^2}{r^2}) - \frac{k}{2\pi} \ln(r/a)$$

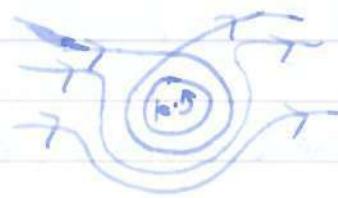
$$k=0$$



high P

$$0 < k < 4\pi a U$$

$$4\pi a U < k$$



low P

Note : For steady potential flow, Bernoulli works everywhere, not just along streamlines.

$$\text{So } p_{\infty} + \frac{1}{2} \rho U^2 = p + \frac{1}{2} \rho \left( \frac{k}{2\pi a} - 2U \sin \theta \right)^2 \text{ on the surface of the circle.}$$

$$\Rightarrow p = p_{\infty} + \frac{1}{2} \rho U^2 - \frac{\rho k^2}{8\pi^2 a^2} + \frac{\rho k}{\pi a} \sin \theta - 2\rho U \sin^2 \theta$$

Symmetric fore and aft, so that there is no net force in the flow-direction.

29/02/12

## Fluid Dynamics (12)

Force in the  $y$  direction:

$$F_y = - \int_0^{2\pi} p \sin \theta \, a \, d\theta = - \int_0^{2\pi} \frac{\rho U k}{\pi a} \sin^2 \theta \times d\theta$$

$$= - \rho U k$$



In general, in "magnum force" resulting from the interaction between a flow  $U$  and a vortex  $K$  is

$$F = \rho U \times K$$

Aerofoils generate circulation around the wing by having a sharp trailing edge, controlling separation.

Kutta Condition - the circulation is just sufficient to cause separation at the leading edge.



Pressure Field in (time dependent) potential flows

Euler E equation:

$$\rho \left( \frac{\partial u}{\partial t} + \nabla \left( \frac{1}{2} |u|^2 \right) - u \times \omega \right) = - \nabla p - \nabla \chi$$

$$\text{If } u = \nabla \varphi, \omega = \nabla \times \nabla \varphi = 0$$

and we have

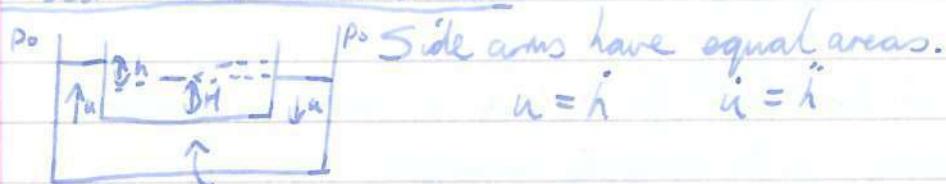
$$\nabla \left[ \rho \frac{\partial \varphi}{\partial t} + \frac{1}{2} \rho |u|^2 + p + \chi \right] = 0$$

$$\Rightarrow \rho \frac{\partial \varphi}{\partial t} + \frac{1}{2} \rho |u|^2 + p + \chi = f(t), \text{ independent of position}$$

In particular, if the flow is steady, then

$$H = \frac{1}{2} \rho |u|^2 + p + \chi \text{ is constant throughout the fluid domain.}$$

Oscillations in a manometer



Large cross section  $\Rightarrow$  Slow flow  $\Rightarrow \varphi \approx \text{constant} = 0$  WLOG

$$\text{LHS } \varphi \approx u y, \frac{\partial \varphi}{\partial t} = i y = h \ddot{y}$$

$$\text{RHS } \varphi \approx -u y, \frac{\partial \varphi}{\partial t} = -i y = -h \ddot{y}$$

Consider pressure on the two free surfaces.

$$\rho h \ddot{y} + \frac{1}{2} \rho h^2 + p_0 + \rho g h = -\rho h \ddot{y} + \frac{1}{2} \rho h^2 + p_0 - \rho g h$$

$$\Rightarrow 2\rho h \ddot{y} + 2\rho g h = 0 \Rightarrow \ddot{y} + \frac{g}{h} h = 0$$

This is the equation for SHM with frequency  $\sqrt{\frac{g}{k}}$

### Oscillations of a bubble

Consider the flow of fluid external to a bubble of radius  $a(t)$ .

$$\nabla^2 \phi = 0 \quad (r > a), \quad \phi \rightarrow 0 \quad (r \rightarrow \infty)$$

$$\frac{\partial \phi}{\partial r} = \dot{a} \quad (r = a)$$

$$\text{Then } \phi = \frac{A}{r} \Rightarrow -\frac{A}{r^2} = \ddot{a} \Rightarrow \phi = -\frac{a^2 \ddot{a}}{r}$$

Ignore gravity and compare a point on the surface of the bubble with the far field

$$\text{Note } \left[ \frac{\partial \phi}{\partial r} \right]_{r=a} = -\frac{a^2 \ddot{a} + 2a\dot{a}^2}{r^2} \Big|_{r=a} = -(a\ddot{a} + 2\dot{a}^2)$$

$$\Rightarrow -\bar{p}(a\ddot{a} + 2\dot{a}^2) + \frac{1}{2}\rho a^2 \dot{a}^2 + p = p_{\infty}(t)$$

$$\rho(a\ddot{a} + 2\dot{a}^2) = p(a, t) - p_{\infty}(t)$$

### Small Oscillations of gas bubbles

$$a = a_0 + \eta(t), \quad \eta \ll a$$

$$p(a_0 + \eta, t) + \frac{1}{2}\rho a^2 \dot{a}^2 = p(a_0, t) + \eta \frac{\partial p}{\partial a} \Big|_{a=a_0} + \frac{1}{2}\rho a^2 \dot{a}^2 - p_{\infty}$$

If oscillations are adiabatic, then  $\rho V^\gamma = \text{constant}$  (where  $\gamma$  is the ratio of specific heats)  $\Rightarrow \frac{\partial p}{\partial a} = -\gamma \frac{\partial V}{\partial a} = -3\gamma \frac{V}{a^2} = -3\gamma \frac{V}{a_0^2} \cdot \left(\frac{a_0}{a}\right)^2$

$$\text{Therefore } p(a_0 + \eta) \approx -3\gamma p_{\infty} \frac{\eta}{a_0^2}$$

$$\Rightarrow \eta \ddot{a} + \frac{3\gamma p_{\infty}}{a_0^2} \eta = 0$$

$$\text{SHM with frequency } \omega = \sqrt{\frac{3\gamma p_{\infty}}{a_0^2}}$$

Frequency  $\approx 2 \times 10^5 \text{ Hz}$  for a 1 mm bubble in water.

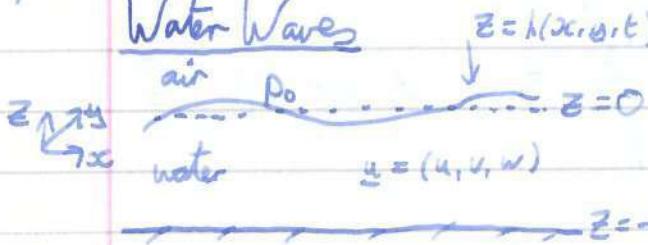
High frequency, so adiabatic assumption is valid.



35/02/12

## Fluid Dynamics (B)

### Water Waves



Assume that the water is incompressible, and motion is started from rest. Then the flow is irrotational,  $\omega = \nabla \times \mathbf{u}$ , and  $\nabla^2 \phi = 0$  ( $-H < z < 0$ )

Kinematic Boundary conditions:

$$\text{At the rigid base } w = \frac{\partial \phi}{\partial z} = 0 \quad (z = -H) \quad (1)$$

$$\begin{aligned} \text{At the free surface } & \frac{dh}{dt} = \frac{\partial \phi}{\partial z} + u \frac{\partial \phi}{\partial x} + v \frac{\partial \phi}{\partial y} = w \quad (z = h) \\ & = \frac{\partial h}{\partial t} + \frac{\partial u}{\partial x} \frac{\partial h}{\partial x} + \frac{\partial v}{\partial y} \frac{\partial h}{\partial y} = \frac{\partial h}{\partial z} \quad (2) \end{aligned}$$

$$\text{Dynamic Boundary Conditions: } p = p_0 \quad (z = h) \quad (3)$$

The expression for pressure in time-dependent potential flow:

$$p + \frac{1}{2} \rho |\nabla \phi|^2 + p_0 + \rho g h = f(t) \quad (z = h) \quad (4)$$

Linearization Assume  $h \ll H$ ,  $\frac{\partial h}{\partial x} \ll 1$   
wave height  $\ll$  wave length  $\lambda$

Also assume that  $|w|$  is small. In consequence of these conditions:

a) Ignore terms of quadratic order or higher in disturbance quantities.

b) Use Taylor expansions to write (for example)

$$\frac{\partial \phi}{\partial z} |_{z=h} = \frac{\partial \phi}{\partial z} |_{z=0} + h \frac{\partial^2 \phi}{\partial z^2} |_{z=0} + \dots \quad \rightarrow \text{quadratic}$$

### Linear Water Waves

$$\nabla^2 \phi = 0 \quad (-H < z < 0) \quad (1)$$

$$\frac{\partial \phi}{\partial z} = 0 \quad (z = -H) \quad (2)$$

$$\frac{\partial h}{\partial t} = \frac{\partial \phi}{\partial z} \quad (z = 0) \quad (3)$$

$$p \frac{\partial}{\partial t} + p_0 + \rho g h = f(t) \quad (z = 0) \quad (4,5)$$

Look for 2D solutions ( $3D$  is not much harder) in the form of a single Fourier Mode (travelling sine wave)

$$h = h(x, t) = h_0 e^{i(kx - \omega t)}$$

( $\omega$  here is the angular frequency, and  $k$  is the horizontal wave number)

$$\phi = \phi(x, z, t) = \hat{\phi}(z) e^{i(kx - \omega t)}$$

Substitute into Laplace's equation:

$$k^2 \phi + \phi'' = 0 \quad (1) \quad \phi' = 0 \text{ at } z = -H \quad (2)$$

$$\Rightarrow \phi = \phi_0 \cosh k(z+H)$$

Boundary conditions at  $z=0$  give

$$-i\omega \phi_0 = k \phi_0 \sinh kH \quad (3)$$

$$-i\omega \phi_0 \cosh kH + g \phi_0 = 0 \quad (4,5)$$

Combining the last two equations:

(Note: these are homogeneous eigenvalue equations for  $\omega(k)$ . They do not determine amplitudes  $\phi_0, \phi_0$ )

$$-\omega^2 \phi_0 \cosh kH + g k \phi_0 \sinh kH = 0$$

$$\Rightarrow [\omega^2 = gk \tanh kH] \quad \text{Dispersion relation}$$

Note that waves of different wavenumbers  $k$  have different frequencies  $\omega$ .

Phase Speed - the speed at which crests travel, for example

$$c = \frac{\omega}{k} \quad [h \propto e^{i(kx - \omega t)} = e^{i(kx - \omega t)}]$$

$$\text{So } c^2 = \frac{\omega^2}{k^2} = \frac{g}{k} \tanh kH$$

and  $c$  depends on wavelength.

Short waves (deep water)

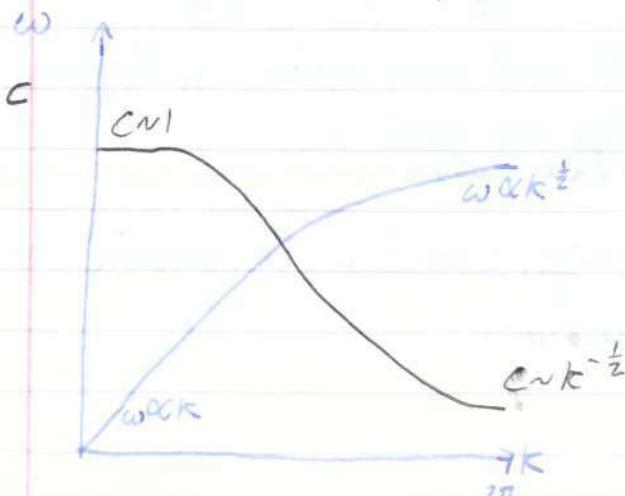
$$\text{wavelength } \lambda = \frac{2\pi}{k} \ll H \Rightarrow kh \gg 1$$

$$\tanh kh \approx 1 \Rightarrow \omega \approx \sqrt{gk}, c \approx \sqrt{\frac{g}{k}}$$

Long waves (shallow water)

$$\lambda \gg H, kh \ll 1, \tanh kh \approx kh$$

$$\Rightarrow \omega \approx \sqrt{gh/k}, c \approx \sqrt{gh}, \text{ constant}$$



07/03/12

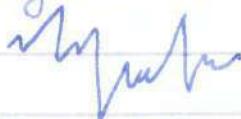
## Fluid Dynamics (14)

### Water Waves

$$\omega^2 = gk \tanh kH$$

Note :

- Unlike light and sound, water waves are dispersive; waves of different lengths travel at different speeds.



Distant Storm

Smooth waves arrive at the beach, long waves first.

- Shallow-water waves (long waves) are approximately non-dispersive

e.g. Tsunamis :  $H \approx 4\text{km}$ ,  $\lambda \approx 500\text{km}$ , shallow water waves

The constant wave speed  $c \approx \sqrt{gH} \approx \sqrt{\frac{10 \times 4 \times 10^3}{m}} = 200\text{ m s}^{-1}$

This is roughly the speed of a commercial airplane.

- \*\* - The group velocity (velocity of energy propagation) is  $C_g = \frac{\partial \omega}{\partial k}$

(See Part II Waves, or Asymptotic methods for proofs.)

So for deep water waves

$$\omega \approx \sqrt{gk}, C_g \approx \sqrt{\frac{g}{k}}, C_g \approx \frac{1}{2} \sqrt{\frac{g}{k}} = \frac{1}{2}c$$

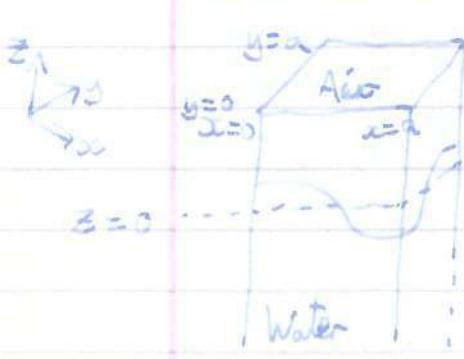
The difference between the phase speed and the group velocity gives rise to the pattern of ship waves. For example :



\* \*

Exercise Re-work the analysis specifically for deep water waves using the boundary condition  $\phi \rightarrow 0$  as  $z \rightarrow -\infty$   
(Strictly,  $\frac{\partial \phi}{\partial z} \rightarrow 0$  as  $z \rightarrow -\infty$ )

Broad Waves e.g. in a square cylinder



Linearised equations :

$$\begin{aligned} \nabla^2 \phi &= 0 & (z < 0) \\ \frac{\partial \phi}{\partial z} &\rightarrow 0 & (z \rightarrow -\infty) \\ \frac{\partial \phi}{\partial x} &\approx 0 & (x = 0, a) \\ \frac{\partial h}{\partial x} &= \frac{\partial y}{\partial x} & (y = 0, a) \\ \frac{\partial z}{\partial x} &= \frac{\partial y}{\partial z} & (z = 0) \\ \Delta \frac{\partial z}{\partial x} + \rho g h &= f(t) & (z = 0) \end{aligned}$$

Separation of Variables :

$$\varphi = \varphi_0 \cos\left(\frac{n\pi x}{a}\right) \cos\left(\frac{m\pi y}{a}\right) e^{kz} e^{i\omega t}$$

$$h = h_0 \cos\left(\frac{n\pi x}{a}\right) \cos\left(\frac{m\pi y}{a}\right) e^{i\omega t}$$

$$k^2 = \frac{\pi^2}{a^2} (n^2 + m^2), \quad \omega^2 = gk \quad (\text{as before})$$

The only difference is that  $k$  is quantized by the side walls. These are standing waves.

Rayleigh - Taylor Instability

The solution (in water) is exactly as for water waves but with  $g \mapsto -g$ .  
Therefore  $\omega^2 = -gk$ ,  $\omega = i\sqrt{gt} \Rightarrow h \propto A e^{Bt^{1/2}} + B e^{-Bt^{1/2}}$

Random Perturbation (initial condition) will have  $A \neq 0$ , so the disturbance will grow in time.

Rayleigh Taylor Instability with surface tension

Consider water waves first.  $\rho = \rho_0 + \gamma * \text{curvature}$

$$\rho = \rho_0 + \gamma * \text{curvature} \approx \rho_0 - \gamma \frac{\partial^2 h}{\partial x^2} \quad \text{surface tension}$$

if  $\frac{\partial h}{\partial x} \ll 1$ . So the linearised dynamic boundary condition becomes

$$\rho \frac{\partial h}{\partial x} + \rho gh - \gamma \frac{\partial^2 h}{\partial x^2} = f(x) \quad (\beta=0)$$

$$\text{In deep water: } \omega^2 = gh + \frac{\gamma}{\rho} h^2$$

This is the dispersion relation for capillary-gravity waves.

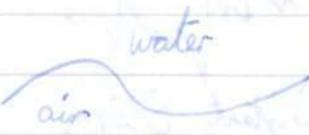
For the Rayleigh Taylor instability,  $\omega^2 = \frac{\gamma}{\rho} k^3 - gk$ .

Instability if  $\omega^2 < 0 \Rightarrow k^2 < \frac{\rho g}{\gamma}$

$$\Rightarrow (n^2 + m^2) \pi^2 < \frac{\rho g a^2}{\gamma}$$

Most unstable if  $n=1, m=0$  (say)

$$\Rightarrow a^2 > \frac{\gamma \pi^2}{\rho g}, \quad a > \sqrt{\frac{\gamma \pi^2}{\rho g}}$$



12/03/12

## Fluid Dynamics (15)

### Fluid Dynamics in a Rotating Frame

The Lagrangian (particle) acceleration in a rotating frame of reference is  $\frac{D\mathbf{u}}{Dt} + 2\Omega \times \mathbf{u} + \Omega \times (\Omega \times \mathbf{x})$  (See Dynamics)

$$\text{So } \rho \left( \frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + 2\Omega \times \mathbf{u} \right) = -\nabla p - \rho \Omega \times (\Omega \times \mathbf{x}) + \rho g$$

where  $\Omega$  is the rotation rate of the frame.



Note that  $\rho \Omega \times (\Omega \times \mathbf{x})$  is a constant (intive) and conservative body force. So the centrifugal force just adds to the hydrostatic pressure, so globally, the sea surface, for example, is not spherical.

For Earth,  $\Omega \approx \frac{2\pi}{10^5} \text{ s}^{-1}$ , Largest scale  $\approx 10^4 \text{ km}$   
So  $\frac{\rho \Omega \times (\Omega \times \mathbf{x})}{\rho g} \leq \frac{(2\pi)^2 \times 10^{-10}}{10} \approx 4 \times 10^{-3} \ll 1$

So we will ignore the centrifugal term.

Consider motions for which  $|\mathbf{u} \cdot \nabla \mathbf{u}| \ll |\Omega \times \mathbf{u}|$   
 $\Rightarrow |\omega| \ll \Omega$

relative vorticity  $\ll$  planetary vorticity

e.g. an atmospheric weather system,  $u = 10 \text{ ms}^{-1}$ ,  $L \approx 10^3 \text{ km}$   
 $\Rightarrow \frac{|\omega|}{\Omega} = \frac{10/10^6}{2\pi \times 10^5} \approx \frac{1}{2\pi}$ , fairly small.

It is a reasonable approximation to ignore the advection term

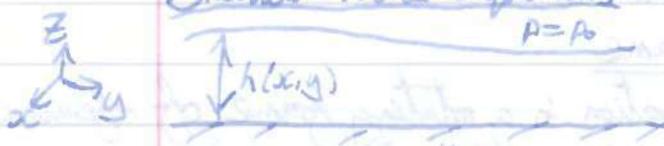
With these two approximations,  $\frac{\partial \mathbf{u}}{\partial t} + 2\Omega \times \mathbf{u} = -\frac{1}{\rho} \nabla p + g$

This is the Euler equation at small Rossby number  $Ro = \frac{|\omega|}{fL} = \frac{1}{f\Omega L}$  in a rotating frame. This approximation relates to

- strong rotation
- low fluid speeds
- large length scales

It is conventional to write  $2\Omega = f$ ,  $f$  called the "planetary vorticity" or the "Coriolis parameter". It is also conventional in this subject to use  $S$  for the relative vorticity  $\nabla \times \mathbf{u}$ .

## Shallow Water Equations



Consider a layer of fluid of depth  $h(x, y)$  with  $p = p_0$  on  $z = h(x, y)$ ,

surface of the ocean or a constant pressure surface defining the top of the atmosphere.

$$\nabla \cdot \underline{u} = 0 \Rightarrow \frac{\partial w}{\partial z} = -\frac{\partial u}{\partial x} - \frac{\partial v}{\partial y}$$

$$\frac{w}{h} \sim \frac{u}{L}, \frac{v}{L}$$

$$\Rightarrow w \sim u \frac{h}{L}, v \frac{h}{L} \Rightarrow w \ll u, v$$

Consider  $\underline{u} = (u, v, 0)$ ,  $\underline{F} = (0, 0, f)$

$$\Rightarrow p \frac{\partial u}{\partial t} - pf v = -\frac{\partial p}{\partial x} \quad (1)$$

$$p \frac{\partial v}{\partial t} + pf u = -\frac{\partial p}{\partial y} \quad (2)$$

$$0 = -\frac{\partial p}{\partial z} - \rho g \quad (3)$$

From (3),  $p = p_0 + \rho g [h(x, y) - z]$  applying  $p = p_0$   
on  $z = h(x, y)$

Horizontal momentum equations

$$(1) \quad p \frac{\partial u}{\partial t} - pf v = -\rho g \frac{\partial h}{\partial x} \quad \text{Note that with our shallow-water}$$

$$(2) \quad p \frac{\partial v}{\partial t} + pf u = -\rho g \frac{\partial h}{\partial y} \quad \text{approximation, pressure is hydrostatic}$$

- Note also that horizontal accelerations are independent of  $z$ .

Initial conditions are usually such that  $u, v$  are also independent of  $z$ .

- Horizontal pressure gradients are proportional to horizontal variations in  $h$ .

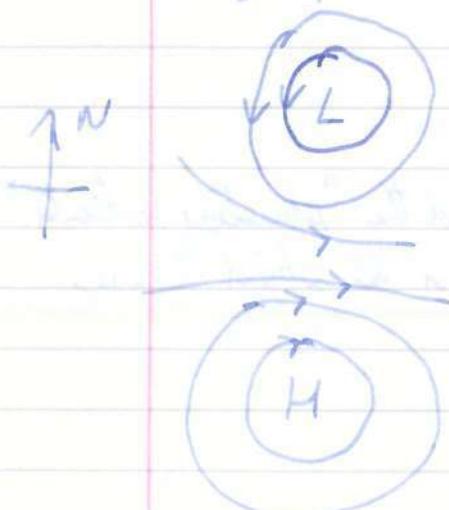
## Geostrophic Balance

$$(1) \quad \text{In steady state, } u = \frac{\partial}{\partial y} \left( -\frac{gh}{f} \right) = \frac{\partial}{\partial y} \left( -\frac{p_0}{f} \right) \quad \text{where } p_0 \text{ is the pressure}$$

$$V = -\frac{\partial}{\partial x} \left( -\frac{gh}{f} \right) = -\frac{\partial}{\partial x} \left( -\frac{p_0}{f} \right) \quad \checkmark \text{ on the base of the}$$

So the 2D streamfunction  $V = -\frac{gh}{f} = -\frac{p_0}{f}$  layer.

Therefore pressure (height) contours are streamlines.



"cyclonic" winds

"anticyclonic" winds

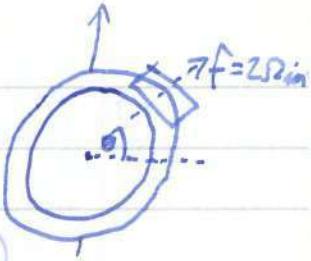
In the Northern Hemisphere, with  
the wind on your back  
Low is on your left  
High is on your right

14/03/12

## Fluid Dynamics (16)

Tangent-plane approximation / f-plane analysis

$$\rightarrow \mathbf{y}(x, y) = (u, v, h) \quad \mathbf{f} = (0, 0, f)$$



### Mass Conservation

Consider a cylinder with horizontal cross-section  $D$

$$\rightarrow u_H \quad \boxed{D} \quad \text{(height)} \quad \frac{d}{dt} \int_D \rho h \, dV = - \int_{\partial D} h \rho u_H \, dS$$

$$\int_D \frac{\partial}{\partial t} (\rho h) \, dV = - \int_D \nabla_H \cdot (\rho h \mathbf{u}_H) \, dV \quad (\nabla_H = \left( \frac{\partial}{\partial x}, \frac{\partial}{\partial y} \right))$$

This is true for arbitrary domains, so  $\frac{\partial h}{\partial t} + \nabla_H \cdot (h \mathbf{u}_H) = 0$

Note that we are still assuming  $\rho = \text{constant}$ .

In Cartesian components  $\frac{\partial h}{\partial t} + \frac{\partial h}{\partial x} (u_x) + \frac{\partial h}{\partial y} (u_y) = 0$ . Note  $\nabla_H \cdot \mathbf{u}_H \neq 0$

### Linearized equations of motion

Suppose  $h = h_0 + \eta(x, y, t)$ ,  $\eta \ll h$

$$\textcircled{1} \quad \text{Then } \frac{\partial \eta}{\partial t} - f v = -g \frac{\partial h_0}{\partial x}, \quad \frac{\partial \eta}{\partial t} + f u = -g \frac{\partial h_0}{\partial y} \quad \Rightarrow \frac{\partial \eta}{\partial t} + f \times \mathbf{v} = -g \frac{\partial h_0}{\partial x} + f \times \mathbf{v} = -g \frac{\partial h_0}{\partial x}$$

$$\frac{\partial^2 \eta}{\partial t^2} + h_0 \left( \frac{\partial^2 \eta}{\partial x^2} + \frac{\partial^2 \eta}{\partial y^2} \right) = 0 \quad \textcircled{2} \quad \frac{\partial^2 \eta}{\partial t^2} + h_0 \nabla \cdot \mathbf{v} = 0$$

(dropping the subscript H)

1. Exercise: For non-rotating, shallow-water waves ( $f = 0$ ),

$$\text{show that } \frac{\partial^2 \eta}{\partial t^2} - g h_0 \left( \frac{\partial^2 \eta}{\partial x^2} + \frac{\partial^2 \eta}{\partial y^2} \right) = 0$$

Wave equation with waves of speed  $c = \sqrt{gh_0}$

2. Eliminate  $\eta$  from  $\textcircled{1}$  by taking the curl:

$$\frac{\partial \mathbf{S}}{\partial t} + (\nabla \cdot \mathbf{u}) \mathbf{f} = 0 \quad \text{where } \mathbf{S} = \left( \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right) \mathbf{e}_z \text{ is the relative vorticity}$$

Substitute for  $\nabla \cdot \mathbf{u}$  from  $\textcircled{2}$

$$\Rightarrow \frac{\partial \mathbf{S}}{\partial t} \left( \mathbf{S} - \frac{\eta}{h_0} \mathbf{f} \right) = 0$$

$\Rightarrow$  Potential Vorticity  $Q = \mathbf{S} - \frac{\eta}{h_0} \mathbf{f}$  is constant in time at each location in space

So  $Q = Q_0(x, y)$  from initial conditions.

Aside: In general (not linearised, or rapidly rotating)

$$\frac{\partial}{\partial t} \left( \mathbf{f} \cdot (\mathbf{S} + \mathbf{f}) \right) = 0 \quad \text{The time}$$

The quantity  $\mathbf{S} + \frac{\eta}{h_0} \mathbf{f}$  is the linearised form of the total potential vorticity,  $\mathbf{f} \cdot (\mathbf{S} + \mathbf{f})$

$\nabla \cdot u$  can be non-zero

$$\textcircled{3} \quad \text{Take the divergence of } \textcircled{1} : \frac{\partial}{\partial t} (\nabla \cdot u) - f \cdot \nabla \times u = -g \nabla^2 \eta$$

Substitute for  $\nabla \cdot u$  from \textcircled{2}

$$-\frac{1}{h_0} \frac{\partial h_0}{\partial t} - f \cdot S = -g \nabla^2 \eta \quad \text{where } S = \nabla \times u \text{ is the relative vorticity}$$

$$\text{Use conservation of potential vorticity to write } S = Q_0 + \frac{\eta}{h_0} f$$

$$\Rightarrow \frac{\partial \eta}{\partial t} - g h_0 \nabla^2 \eta + f^2 \eta = -h_0 f \cdot Q_0, \text{ independent of time}$$

Note that the unforced equation ( $Q_0 = 0$ ) supports waves with  $\eta \propto e^{i\omega t}$

Example Suppose that there is a region of high pressure next to a region of low pressure.

$$h_0 - \eta_0 \quad \underbrace{\dots \dots \dots}_{h_0} \quad \overbrace{\dots \dots \dots}^{h_0 + \eta_0}$$

In the non-rotating case, we get waves

travelling in both directions.

In the rotating case, there is a non-trivial, geo-strophically balanced steady flow. Note that

$$Q_0 = S_0 - \frac{\eta_0}{h_0} f = \mp \frac{\eta_0}{h_0} f \text{ in } x \gtrless 0$$

$$\frac{\partial \eta}{\partial t} - g h_0 \nabla^2 \eta + f^2 \eta = \pm f^2 \eta_0 \text{ in } x \gtrless 0$$

Steady (Geo-strophically balanced) flow has  $\eta = \eta(x)$

$$\eta'' - \frac{1}{R^2} \eta = \mp \frac{1}{R^2} \eta_0$$

where  $R = \frac{f h_0}{g}$ , is the Rossby radius (of deformation).

Solve with  $\eta \geq 0$  as  $x \rightarrow \pm \infty$

$\eta, \eta'$  continuous at  $x = 0$

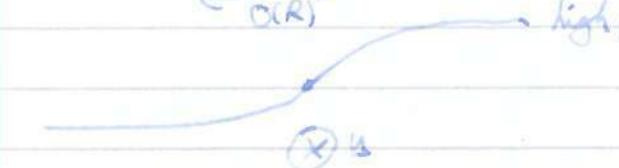
↳ Continuity of  $u$

↳ This gives continuity of pressure.

14/03/12

### Fluid Dynamics ⑬

$$\Rightarrow \eta = \begin{cases} \eta_0 (1 - e^{-\frac{x}{R}}) & x > 0 \\ -\eta_0 (1 - e^{\frac{x}{R}}) & x < 0 \end{cases}$$

$$u = \underbrace{\frac{\partial}{\partial y} \left( -\frac{\partial \eta}{\partial x} \right)}_{O(R)} = 0, \quad v = -\frac{\partial}{\partial x} \left( -\frac{\partial \eta}{\partial x} \right) = \eta_0 \sqrt{\frac{g}{h_0}} e^{-\frac{|x|}{R}}$$


The Rossby radius gives the characteristic length scale for balanced flow in the atmosphere and oceans.

In plan-view:



This is like a section through



